



## Article (refereed) - postprint

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Ward, Susan; Ostle, Nicholas J.; Oakley, Simon; Quirk, Helen; Henrys, Peter A.; Bardgett, Richard D.. 2013. **Warming effects on greenhouse gas fluxes in peatlands are modulated by vegetation composition.** *Ecology Letters*, 16 (10). 1285-1293. <https://doi.org/10.1111/ele.12167>

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**Warming effects on greenhouse gas fluxes in peatlands are modulated by vegetation composition**

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Running title: Plants modulate warming effects on GHG fluxes

Keywords

Warming, greenhouse gas, plant functional group, plant community composition, peatland, CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O, carbon cycle

24

25 Article type: Letter

26

27 Word count:

28 Abstract: 149 words

29 Main text: 4496 words

30 Text box: n/a

31

32 Number of references: 45

33 Number of Figures: 4

34 Number of Tables: 2

35 Number of text boxes: 0

36

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40 Authorship statement.

41 RDB and NJO conceived and designed the experiment, with input into the design from SEW.

42 SEW, SO, HQ performed the study, collected the data, and analysed the samples. SEW and

43 PAH analysed the data, and SEW, RDB, NO wrote the paper, to which all authors

44 contributed with discussions and text.

45

**ABSTRACT**

Understanding the effects of warming on greenhouse gas feedbacks to climate change represents a major global challenge. Most research has focused on direct effects of warming, without considering how concurrent changes in plant communities may alter such effects. Here, we combined vegetation manipulations with warming to investigate their interactive effects on greenhouse gas emissions from peatland. We found that although warming consistently increased respiration, the effect on net ecosystem CO<sub>2</sub> exchange depended on vegetation composition. The greatest increase in CO<sub>2</sub> sink strength after warming was when shrubs were present, and the greatest decrease when graminoids were present. CH<sub>4</sub> was more strongly controlled by vegetation composition than by warming, with largest emissions from graminoid communities. Our results show that plant community composition is a significant modulator of greenhouse gas emissions and their response to warming, and suggest that vegetation change could alter peatland carbon sink strength under future climate change.

## INTRODUCTION

There is growing concern about how biosphere carbon dynamics will respond to expected climate change, with evidence suggesting that atmospheric warming will increase soil respiration and greenhouse gas feedbacks (Bardgett *et al.* 2008; Craine *et al.* 2010). At the same time, terrestrial ecosystems are being subjected to increasing environmental pressures and human demands that are affecting vegetation community composition and diversity globally (Thuiller *et al.* 2005; Stevens *et al.* 2010). Despite widespread recognition that both climate and vegetation change can act independently as drivers of ecosystem carbon dynamics (De Deyn *et al.* 2008; Dorrepaal *et al.* 2009), we know little about the potential role in the carbon cycle of interactions between them (Bardgett *et al.* 2013). Indeed, experiments that explore the independent and interactive effects of abiotic and biotic factors as controls over ecosystem functioning are few (Hooper *et al.* 2005; Kardol *et al.* 2010), despite the suggestion that the magnitude of effects of vegetation change on ecosystem processes can be comparable to that of environmental change (Hooper *et al.* 2012; Tilman *et al.* 2012).

Carbon rich peatlands provide an ideal model system in which to examine the influence of warming and vegetation change on ecosystem greenhouse gas emissions; they have a relatively simple plant community structure, are recognised as important global sinks and sources of the greenhouse gases CO<sub>2</sub> and CH<sub>4</sub> respectively, and are vulnerable to land use and climate change (Dise 2009). Climate change models predict that northern latitude peatlands will be subjected to higher temperatures with longer growing seasons (IPCC 2007), and that this change will be accompanied by an increase in vascular plants at the expense of bryophytes and lichens (Walker *et al.* 2006; Gallego-Sala & Prentice 2013). Recent work has shown that experimental warming can significantly increase rates of peatland ecosystem respiration (Dorrepaal *et al.* 2009; Briones *et al.* 2010), and that drought can induce carbon

loss *via* changes in soil enzyme activity (Fenner & Freeman 2011). Other peatland studies suggest that there are key differences in the ecophysiological traits of dominant plant functional groups, that have a strong regulatory role in ecosystem carbon dynamics (Ward *et al.* 2012). Despite this, it is not known whether changes in plant community structure, and the presence or absence of dominant peatland plant functional groups (*i.e.*, shrubs, graminoids and bryophytes), will modify the impact of warming on ecosystem greenhouse gas fluxes. This represents a serious knowledge gap, given that most plant communities globally are subject to both vegetation and climate change, but their combined impact on greenhouse gas fluxes is not known.

To redress this gap in knowledge, we established a unique field experiment in spring 2008 with the aim of examining the independent and interactive effects of warming and plant functional composition on greenhouse gas exchange in a peatland ecosystem. We used a plant removal approach to manipulate vegetation composition (Diaz *et al.* 2003; Wardle & Zackrisson 2005) from an area of ombrotrophic blanket bog in northern England. Vegetation manipulations included removal of all possible combinations of the three dominant plant functional groups, namely ericoid shrubs, graminoids (sedges), and bryophytes/lichens. Warming was induced passively on half of the experimental plots, using randomly allocated hexagonal open-top chambers (OTCs) (Marion *et al.* 1997) which increased air temperatures by approximately 1°C over the mid-day period. We present results from field measurements of greenhouse gas fluxes, namely net ecosystem exchange (NEE) of CO<sub>2</sub>, ecosystem respiration, CH<sub>4</sub> and N<sub>2</sub>O fluxes for all times of year spanning two growing seasons. We show that, although rates of ecosystem respiration were consistently increased by warming across all vegetation types, the effect of warming on NEE, once differences in photosynthetic uptake of CO<sub>2</sub> were taken into account, was dependent on plant community composition.

110 More specifically, the greatest increase in CO<sub>2</sub> sink strength after warming was observed  
111 when shrubs only (dominated by *Calluna vulgaris*) were present. Also, warming reduced  
112 mean CO<sub>2</sub> sink strength in the presence of graminoids, and increased CO<sub>2</sub> sink strength when  
113 graminoids were absent. In addition, we found that the efflux of CH<sub>4</sub> was more strongly  
114 controlled by plant community composition than by warming, with largest emissions coming  
115 from sedge (*Eriophorum vaginatum*) dominated communities. Taken together, these findings  
116 highlight the importance of plant community composition as a driver of carbon cycling  
117 processes, and show that plant community composition can modulate the effects of warming  
118 on net ecosystem exchange of CO<sub>2</sub>.

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## MATERIALS AND METHODS

### Study site.

The study site was situated on an area of ombrotrophic blanket bog within the Moor House National Nature Reserve in northern England (54°65' N, 2°45' W). The site altitude was 550m, the mean annual temperature is 5.8°C, and the mean annual precipitation 2048mm (UK Environmental Change Network, [www.data.ecn.ac.uk](http://www.data.ecn.ac.uk)). The mean depth of peat at the site was 1.17m ( $\pm 0.01$ ), and the mean pH 4.07 ( $\pm 0.01$ ). Abiotic conditions, including air and soil temperature, solar radiation, photosynthetically active radiation and rainfall at the site were recorded by the Moor House automated weather station ([www.ecn.ac.uk](http://www.ecn.ac.uk)) (Supporting information, Table S1).

### Vegetation manipulation and climate warming.

Vegetation removals were undertaken by hand, from areas measuring 1.5 x 1.5m, separated by a buffer zone of at least 1m from adjoining plant removal plots. Shoots of shrubs and graminoids were cut back to litter layer level, and all green (photosynthetic) tissues of bryophytes were removed, taking care to minimise disturbance of the soil and remaining vegetation types. Wooden boardwalks were installed on two sides of each removal plot, to allow access to the sampling plots without damage by trampling. Plots were left to settle for a year before sampling to minimise effects of decomposition from roots. The use of this plant removal approach allowed us to measure the effects of plant functional groups *in situ* in their natural environment (Diaz *et al.* 2003). The plant functional group manipulations were from the three dominant vegetation types present: ericoid dwarf-shrubs (S), dominated by *Calluna vulgaris* (L.) Hull; graminoids (G), dominated by the sedge *Eriophorum vaginatum* L.; and bryophytes/lichens (B) dominated by feather mosses (*Hypnum jutlandicum* Holm. &



Warncke; *Pleurozium schreberi* (Brid.) Mitt.) and *Sphagnum* mosses. There were 8 different plant manipulations: a control with all vegetation present, three single groups (S, G or B), three double groups (S&G, S&B, G&B) and a treatment where all above ground vegetation was removed. The experiment site had four blocks, containing randomly arranged warmed and non-warmed replicates of each plant manipulation treatment (n = 64).

Warming was achieved passively using hexagonal OTCs based on the ITEX design (Marion *et al.* 1997), modified for peatland vegetation by the addition of a 20cm high vertical galvanised steel base, on to which the transparent top sections were fixed using cable ties. Each transparent section making up the hexagonal OTC measured 80cm along the bottom edge, 62.5cm along the top edge and 40cm height, to give an internal diameter of 1m<sup>2</sup>, avoiding edge effects. The transparent material was 2mm thick Liteglaze clear acrylic sheet (Ariel plastics, UK), which allows 92% light transmission. The open-topped chamber method offers a robust means to examine effects of warming in remote environments, without the need for a power supply, and has been used frequently in arctic and peatland ecosystems (Walker *et al.* 2006; Dorrepaal *et al.* 2009). This methodology has its limitations, most notably that OTCs can act as a physical barrier to wind (Marion *et al.* 1997), which, in addition to changing temperature, has the potential to alter the width of the boundary layer and hence the concentration of CO<sub>2</sub> surrounding photosynthesising leaves, thereby affecting rates of photosynthetic uptake of carbon. Despite these limitations, the technique provides a valid and useful way of quantitatively comparing the effects of warming between experimental vegetation removal treatments in the field.

The OTCs were fixed in place one month prior to commencement of sampling. Air temperatures at vegetation canopy height were recorded using temperature loggers (Lascar

Electronics, Salisbury, UK). Water table levels were measured from dip-wells made of 1m long perforated PVC pipe, installed in each of the 64 experimental plots. On average, the OTC's increased mean air temperatures by 0.88°C and 0.72°C over the midday period (during the gas sampling period between 11:00 and 14:00 hrs), for the growing and non-growing season respectively. Over 24 hours, the mean increase in temperature was 0.46°C and 0.21°C for the growing and non-growing seasons. We found no evidence of any difference in water table draw-down due to warming ( $F_{1,1476} = 0.2$ ,  $P = 0.87$ ), or due to vegetation type ( $F_{1,1476} = 0.9$ ,  $P = 0.33$ ). For full details of the abiotic conditions during all sampling dates, see Table S1.

#### **Greenhouse gas flux measurements.**

In each sampling plot, a 30 cm diameter, 10cm high gas sampling base ring was fitted in place at 5cm depth, with care taken to minimise disturbance and to avoid severance of large plant roots. Boardwalks installed on two sides of each experimental plot allowed access to the sampling areas without compressing the surrounding peat, which could have created physical movement of gases. Measurements of CO<sub>2</sub> exchange were made over 120-s intervals with a PP systems EGM4 portable IRGA coupled to a customised chamber lid, 30cm diameter and 35cm height (Ward *et al.* 2007). We used the dark and light flux method for ecosystem respiration and net CO<sub>2</sub> flux respectively (Ward *et al.* 2007). Measurements were taken between 11:00 and 14:00 hours from June 2009 to August 2010, at approximately monthly intervals during the growing season, and bi-monthly at other times. For CH<sub>4</sub> and N<sub>2</sub>O, bi-monthly gas samples were collected on closure of the chamber lid and at three additional time points up to 30 minutes closure. Gas samples (10ml) were taken from the chamber headspace using a gas syringe, and injected into evacuated 3ml exetainers (Labco, UK) for storage prior to analysis. Concentrations of CH<sub>4</sub> and N<sub>2</sub>O were analysed by gas

chromatography, using Perkin Elmer Autosystem XL GCs with a flame ionisation detector for CH<sub>4</sub> and electron capture detector for N<sub>2</sub>O. GC detection limits were better than 0.2 ppm for all gases. For each sample, 2.5ml of gas was injected into the GC using an HTA Autosampler. Results were calibrated against certified gas standards, comprising 500ppm CO<sub>2</sub>, 10ppm CH<sub>4</sub> and 1ppm N<sub>2</sub>O (BOC, UK). All fluxes were adjusted for field sampling temperature, headspace volume and chamber area (Holland *et al.* 1999), and calculated by linear regression using all time points sampled (Levy *et al.* 2012).

### **Soil properties**

Peat cores measuring 3cm diameter and 10cm depth were collected from each field plot in July of the final year of gas sampling, in order to gain a measure of microbial biomass and the availability of dissolved organic carbon (DOC) and nitrogen (DON) in the peat. Peat was homogenised and hand sorted to remove any root material, then analysed for microbial biomass C and N using fumigation-extraction, and water extractable DOC and DON using methods described in Ward *et al.* (2007).

### **Statistics.**

Data were checked for normality using residual plots method, and log-transformed where necessary before analysis. The effects of experimental warming and vegetation manipulations, and their interactions, were analysed by repeated measures ANOVA, using SAS Enterprise Guide 4, with sampling date nested within sampling block as random effects. Vegetation effects were analysed as the presence and absence of each of the three plant functional groups (shrubs, graminoids and bryophytes), and effects of vegetation diversity were analysed based on the number of plant functional groups present. After confirming a three way interaction between season, warming and plant functional group, data were

220 analysed as 2 separate models: 1) growing season data; and 2) non growing season data, with  
221 growing season defined as when the mean air temperature is greater than 6°C.

222

## RESULTS

### CO<sub>2</sub>

Our results show that the effect of warming on NEE of CO<sub>2</sub> was modulated by the removal of different plant functional groups in the experimental communities (Fig 1, Table 1). A significant interaction ( $F_{1,744} = 6.4$ ,  $P = 0.0126$ ) between warming and plant functional group removal on NEE was observed during the growing season (*i.e.* when average air temperature was  $> 6^{\circ}\text{C}$ ). More specifically, mean CO<sub>2</sub> sink strength increased by 55% with warming in plots where shrubs were the only plant functional group present, and by 36% when shrubs were present with bryophytes, but without graminoids (Fig. 1). In the presence of graminoids, however, mean CO<sub>2</sub> sink strength was reduced by 20% with warming, whereas in the absence of graminoids, mean CO<sub>2</sub> sink strength was increased by 43% with warming. Vegetation diversity also influenced NEE ( $F_{3,744} = 18.3$ ,  $P < 0.0001$ ), with strongest effects seen when comparing non-vegetated plots with those containing 2 or 3 plant functional groups, but there were no interactions between vegetation diversity and warming (Supporting information, Table S2).

Ecosystem respiration rates were consistently raised by warming across all vegetation treatments (Fig. 2), but there were no detectable interactions of warming with the removal of shrubs, graminoids or bryophytes (Table 1). Across all vegetation removal treatments, warming of  $\sim 1^{\circ}\text{C}$  over the year increased rates of ecosystem respiration in warmed relative to non-warmed treatment plots by a mean of 47% and 49%, during the growing and non-growing seasons respectively ( $F_{1,734} = 49.8$ ,  $P < 0.0001$ ;  $F_{1,227} = 10.1$ ,  $P = 0.002$ ). There were also highly significant effects of shrub, graminoid, and bryophyte removal on ecosystem respiration rates, with strongest effects during the growing season, and interactions observed between graminoids and the other plant functional groups (Table 1). The highest rates of

respiration were measured in the presence of vascular plants, and there was a greater reduction in respiration from the removal of shrubs than from the removal of graminoids. When bryophytes were removed, rates of respiration increased, however this effect was only observed during the growing season (Table 1). Significantly lower rates of respiration were measured for bare plots compared to those with one or more plant functional group present ( $F_{3,734} = 21.1$ ,  $P < 0.0001$ ), but there was no interaction of vegetation diversity with warming (Supporting information, Table S2).

### **CH<sub>4</sub> and N<sub>2</sub>O**

We found that vegetation composition, particularly the presence of graminoids, was a stronger factor than warming in regulating peatland CH<sub>4</sub> fluxes (Fig. 3, Table 1), and that the presence and absence of graminoids and shrubs interacted to affect net CH<sub>4</sub> exchange all year round. Emissions of CH<sub>4</sub> were higher in the presence relative to absence of graminoids, but lower in the presence relative to absence of shrubs. We measured the highest CH<sub>4</sub> fluxes when graminoids (the sedge, *Eriophorum vaginatum*) were present without shrubs and without bryophytes (Fig. 3). Warming effects on CH<sub>4</sub> efflux were only significant during the growing season ( $F_{1,251} = 5.6$ ,  $P = 0.02$ ), but we detected no interactive effect of warming with vegetation removal on ecosystem CH<sub>4</sub> emissions for any of the three plant functional groups (Table 1). Outside the growing season, the peatland was seen to be a small sink for CH<sub>4</sub> in the absence of vegetation, and when shrubs and bryophytes only were present in warmed plots (Fig. 3).

For N<sub>2</sub>O, we found no significant effect of warming either during ( $F_{1,167} = 0.0$ ,  $P = 0.92$ ) or outside the growing season ( $F_{1,167} = 2.5$ ,  $P = 0.12$ ), although there was a trend for a greater N<sub>2</sub>O sink in warmed plots during the non-growing season (Fig. 4, Table 1). During the

growing season we detected an interactive effect between shrubs and bryophytes, whereby the greatest mean sink for N<sub>2</sub>O was measured when shrubs were present and bryophytes had been removed. There were no interactions between warming and vegetation removal, and no significant effect of plant diversity on N<sub>2</sub>O flux.

### **Soil properties**

Warming increased concentrations of DOC ( $F_{1,64} = 6.1$ ,  $P = 0.02$ ) and DON ( $F_{1,64} = 7.0$ ,  $P = 0.01$ ) in soil solution by 13% and 15% respectively (Table 2). Vegetation change was found to have a stronger effect on DOC and DON than warming, with the removal of shrubs increasing concentrations of DOC ( $F_{1,64} = 22.4$ ,  $P < 0.0001$ ) and DON ( $F_{1,64} = 21.0$ ,  $P < 0.0001$ ) by 21%. In contrast, the graminoid or bryophyte removal had no detectable effect on DOC or DON, and no interactions between warming and vegetation change were detected (Supporting information, Table S3). Microbial biomass C and N did not respond to warming, although microbial N was affected by vegetation change: microbial biomass N was greatest when both shrubs and bryophytes were removed (Supporting information, Table S3), and microbial C:N ratio was 14% lower when shrubs were removed ( $F_{1,64} = 5.4$ ,  $P = 0.02$ ).

## DISCUSSION

It has long been recognised that carbon cycling processes in peatlands are highly sensitive to changes in climate (Dise 2009; Dorrepaal *et al.* 2009), and there is growing evidence that climate driven vegetation change in peatland and high latitude ecosystems is leading to an increase in vascular plants at the expense of bryophytes (Walker *et al.* 2006; Gallego-Sala & Prentice 2013). However, despite these concurrent changes in climate and vegetation, their interactive effects on greenhouse gas fluxes are virtually unknown. We, therefore, set out to examine the independent and interactive effects of warming and plant functional composition on greenhouse gas exchange in a peatland ecosystem, using a unique field plant manipulation and warming experiment. Our findings provide the first evidence that the response of peatland greenhouse gas exchange to warming is both modulated and strongly controlled by plant community composition.

Our results show that removal of different plant functional groups in the experimental communities modulated the effects of warming on NEE of CO<sub>2</sub>. In particular, we found that, during the growing season, a significantly greater increase in net CO<sub>2</sub> sink strength with warming was seen in the presence of shrubs when graminoids were absent, whereas warming had the opposite effect in the presence of graminoids. As the main terrestrial exchange of carbon from peatlands is as CO<sub>2</sub> (Roulet *et al.* 2007), quantifying NEE of CO<sub>2</sub> allows us to get a measure of the net ecosystem carbon balance of the system, and how this is affected by warming and vegetation community composition. The clear interactive effect of warming and plant functional group removal on NEE during the growing season (*i.e.* when average air temperature was > 6°C), suggests that responses were dependent on feedbacks from actively growing plants, supporting the idea that the composition of actively growing peatland vegetation is a key modulator of the response of ecosystem CO<sub>2</sub> fluxes to climate change. In



contrast, although ecosystem respiration rates were consistently raised by warming across all vegetation treatments, no such interaction of warming with vegetation composition was detected. Given the similarity of the respiration responses of different plant functional groups to warming, we propose that observed differences in NEE are largely attributable to differences in photosynthetic CO<sub>2</sub> uptake, with shrubs growing alone, or shrubs with bryophytes, showing the greatest increase in photosynthesis relative to respiration and hence, increased net CO<sub>2</sub> sink strength, with warming. In contrast, in the presence of graminoids, warming led to a greater increase in rates of respiration relative to photosynthesis, leading to a reduction in net CO<sub>2</sub> sink strength. Differences in rates of assimilation of CO<sub>2</sub> and translocation of new photosynthates below-ground have previously been observed among dominant peatland plant functional groups (Ward *et al.* 2012), with vascular plants (shrubs and graminoids) showing greater rates of CO<sub>2</sub> assimilation and transfer relative to bryophytes. This significant positive effect of warming on photosynthetic drawdown of CO<sub>2</sub> by shrubs is likely to be a consequence of their characteristic ecophysiological traits related to resource acquisition, including associations with ericoid mycorrhizal fungi (Read *et al.* 2004), and canopy height and bushy growth habit, which makes them better placed to intercept light and to shade vegetation beneath their canopy.

Another explanation for the differences in warming response of NEE across plant functional groups might be associated shifts in microbial communities in the peat, which could ultimately affect the balance between CO<sub>2</sub> uptake and release under warming (Bardgett *et al.* 2008). It is known that shrubs and graminoids in peatlands differ in the rate that they allocate photosynthetic carbon below-ground (Ward *et al.* 2009; Ward *et al.* 2012), and such differences in allocation are likely to affect the quality and quantity of exudates released from roots, to mycorrhizal fungi, and ultimately to soil, thereby affecting the composition and

activity of microbial communities (De Deyn *et al.* 2008; Bardgett *et al.* 2013). Also, observed differences in the photosynthetic response of plant functional groups to warming are likely to have altered carbon flux to roots and rates of root exudation, thereby further contributing to shifts in the composition and activity of microbial communities across vegetation treatments, and potentially explaining differential responses of NEE to warming. We did not measure soil microbial community structure in this study, but we did find, albeit at one sample date, that microbial C:N was significantly affected by shrub removal, which could be indicative of a change in microbial communities. This is perhaps due to the high concentrations of phenolic compounds (Hattenschwiler & Vitousek 2000; Freeman *et al.* 2001) and the presence of mycorrhizal fungi (Read *et al.* 2004; Orwin *et al.* 2011) associated with ericoid shrubs. More studies are clearly needed to unravel the mechanisms by which differences in vegetation modulate responses of NEE to warming, including studies on the potential role of shifts in microbial communities as determinants of the response of NEE to warming.

The mean increase in rates of ecosystem respiration in response to ~1°C warming, of 47-49%, is consistent with other studies of warming effects in peatlands, observed in the field (Dorrepaal *et al.* 2009) and laboratory (Kim *et al.* 2012). As with our findings for NEE of CO<sub>2</sub>, the greatest effects of vegetation composition were observed during the growing season, with the highest respiration rates being measured when vascular plants (*i.e.*, shrubs and graminoids) were present in the plant community. We also observed warming effects on concentrations of DOC and DON in soil solution, which were found to be higher in soils from the warmed than unwarmed plots at the peak of the growing season, which is likely indicative of an increase in microbial activity in response to warming. Although the effects of warming and vegetation composition on ecosystem respiration were found to be

independent, our findings do highlight the importance of both warming and actively growing vegetation in controlling the release of CO<sub>2</sub> to the atmosphere by respiration.

Peatlands are a globally important source of CH<sub>4</sub> (Baird *et al.* 2009) and previous work has shown that both warming (van Winden *et al.* 2012) and vegetation (Levy *et al.* 2012; Gray *et al.* 2013) can have a measureable effects on ecosystem CH<sub>4</sub> emissions. Our study provides the first *in situ* field experimental evidence that vegetation composition is a stronger factor than ~ 1°C warming in regulating peatland CH<sub>4</sub> fluxes. As with our findings for CO<sub>2</sub>, the effects of vegetation community composition were stronger during the growing season than for the rest of the year, which again, highlights the key role that actively growing vegetation can play in controlling GHG exchange. The relatively greater levels of CH<sub>4</sub> efflux when the graminoid *Eriophorum vaginatum* was present, may be explained by the recognised functional traits of this wetland sedge, namely the presence of aerenchymous tissues which act as a conduit for CH<sub>4</sub> from the catotelm (Strack *et al.* 2006; Green & Baird 2012). In addition, differences in the quality and quantity of root exudates entering the soil from contrasting plant functional groups (De Deyn *et al.* 2008) are likely to affect the activity of methanogenic bacteria in the peat, as well as respiration processes. Sedges in particular have been associated with enhanced CH<sub>4</sub> production due to an increased supply of available substrates, particularly acetate, to methanogens (Bellisario *et al.* 1999; Hornibrook 2009; Lai 2009), providing an additional explanation for the increased CH<sub>4</sub> emissions we observed in the presence of graminoids. Interestingly, we observed that the system was a small sink for CH<sub>4</sub> in the absence of vegetation, and also when shrubs and bryophytes only were present in warmed plots outside the growing season. This implies that, even out of the growing season, the presence of vegetation is still exerting controls on CH<sub>4</sub> emissions, through either changes in microbial activity, or differences in physical conditions. Although peatlands are associated

with greater CH<sub>4</sub> productivity than consumption, there is some evidence of CH<sub>4</sub> consumption in peat, particularly by methylocystis-related species (Kolb & Horn 2012).

Whereas previous peatland observations of high CH<sub>4</sub> emissions from sedge dominated communities come from contrasting physical habitats (Strack *et al.* 2006; McNamara *et al.* 2008), our plant manipulation approach allowed us to compare plant functional groups in the same habitat, providing new evidence of the importance of vegetation composition in controlling CH<sub>4</sub> emissions. Although warming did increase the mean CH<sub>4</sub> efflux for all vegetation manipulation treatments by 90% during the growing season, these effects were much weaker than those observed due to vegetation composition, and we found no statistical evidence that warming effects differed between vegetation types.

Atmospheric exchange of N<sub>2</sub>O, the third greenhouse gas measured in this study, was relatively low, as would be expected in nutrient poor ecosystems such as peatlands (Reay *et al.* 2012). These small fluxes, which varied between net emissions and net uptake (Fig. 4), are typical of northern ombrotrophic peatlands (Drewer *et al.* 2010). Despite this, vegetation composition was found to impact on sink strength for N<sub>2</sub>O during the growing season, being greatest when shrubs were present and bryophytes had been removed. Unlike for other greenhouse gases, however, we detected no response of warming on N<sub>2</sub>O, aside a weak increase in N<sub>2</sub>O sink strength, suggesting that climate warming is unlikely to affect the atmospheric exchange of N<sub>2</sub>O in peatland in this N poor blanket peatland. In contrast, studies of N<sub>2</sub>O emissions from peatlands which are more nutrient-rich (Martikainen *et al.* 1993), or which have patches of bare soil with high nitrate content due to cryoturbation (Repo *et al.* 2009), have shown that climate warming can have powerful effects on peat N<sub>2</sub>O fluxes.

In conclusion, our findings provide evidence from a unique field manipulation experiment that warming effects on greenhouse gas exchange in peatland are modulated by changes in plant community composition, with the greatest increase in net CO<sub>2</sub> sink strength with warming occurring when shrubs were present and graminoids were absent. A change in the rate of greenhouse gas exchange with the atmosphere, brought about by increased domination by vascular plants as peatlands warm, has the potential to feedback to global climate change by exacerbating radiative forcing. Furthermore, the observed interaction of climate warming with vegetation change could accelerate these feedbacks in peatland systems containing large stocks of globally important carbon (Gallego-Sala & Prentice 2013). Whilst the mechanisms that underlie our findings require further exploration, our results indicate that changes in vegetation community composition can act as a strong determinant of climate change effects on northern peatland carbon cycling. As such, these results highlight the importance of considering biotic as well as abiotic climate induced changes when predicting the future greenhouse gas sink/source strength of peatland ecosystems.

430 **ACKNOWLEDGEMENTS**

431 This research was supported by a Natural Environment Research Council (NERC) EHFI  
432 grant (NE/E011594/1) awarded to R D Bardgett, as lead investigator, and N J Ostle. We  
433 thank colleagues from CEH Lancaster and Lancaster University for help in the field. We also  
434 thank Natural England and the Environmental Change Network, CEH Lancaster, for site  
435 access and weather station data. We are grateful to three anonymous referees for their  
436 valuable comments on an earlier version of this manuscript.

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607 **SUPPORTING INFORMATION**

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609 Wiley Online Library ([www.ecologyletters.com](http://www.ecologyletters.com)).

610 **Table S1.** Abiotic conditions for all sampling dates.

611 **Table S2.** Effects of vegetation diversity on CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O.

612 **Table S3.** Statistical analysis for DOC, DON and microbial biomass C and N.

613

614 **TABLES**

615 **Table 1.** Statistical analysis for the effects of, and interactions between, warming and the  
616 presence/absence of plant functional groups on CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O fluxes by seasons: a)  
617 growing season May to September, b) non-growing season October to April.

618

619 Table 1

Source of variation	a) Growing season (May – Sept)			b) Non-growing season (Oct – April)		
	df	f	p	df	f	p
<b><i>Net ecosystem CO<sub>2</sub> exchange (mg m<sup>-2</sup> h<sup>-1</sup>)</i></b>			(n = 744)			(n = 229)
Warming	1	0.0	0.86	1	3.1	0.08
Shrub presence/absence	1	151.4	<0.0001	1	21.1	<0.0001
Graminoid presence/absence	1	17.1	<0.0001	1	2.8	0.10
Bryophyte presence/absence	1	0.1	0.82	1	1.6	0.21
Warmed x shrub	1	0.5	0.48	1	2.5	0.12
Warmed x graminoid	1	8.0	0.005	1	0.0	0.97
Warmed x bryophyte	1	0.4	0.52	1	1.5	0.22
Shrub x graminoid	1	13.4	0.0004	1	0.0	0.95
Shrub x bryophyte	1	2.7	0.10	1	1.7	0.20
Graminoid x bryophyte	1	2.9	0.09	1	1.9	0.18
Warmed x shrub x graminoid	1	6.4	0.01	1	0.1	0.81
Warmed x shrub x bryophyte	1	0.0	0.87	1	1.4	0.25
Warmed x graminoid x bryophyte	1	0.2	0.67	1	0.1	0.77
<b><i>Ecosystem respiration (mg m<sup>-2</sup> h<sup>-1</sup>)</i></b>			(n = 734)			(n = 227)
Warming	1	49.8	<0.0001	1	10.1	0.002
Shrub presence/absence	1	164.7	<0.0001	1	22.6	<0.0001
Graminoid presence/absence	1	32.8	<0.0001	1	6.1	0.016
Bryophyte presence/absence	1	5.4	0.022	1	0.4	0.55
Warmed x shrub	1	1.0	0.31	1	0.5	0.51
Warmed x graminoid	1	0.2	0.63	1	0.2	0.69
Warmed x bryophyte	1	0.5	0.50	1	0.1	0.72
Shrub x graminoid	1	14.3	0.0002	1	2.4	0.12
Shrub x bryophyte	1	0.0	0.90	1	1.3	0.26
Graminoid x bryophyte	1	19.3	<0.0001	1	2.0	0.16
(no significant 3 way interactions)						
<b><i>CH<sub>4</sub> (mg m<sup>-2</sup> h<sup>-1</sup>)</i></b>			(n = 251)			(n = 218)
Warming	1	5.6	0.02	1	0.3	0.58
Shrub presence/absence	1	9.9	0.002	1	1.5	0.22
Graminoid presence/absence	1	10.1	0.002	1	15.9	0.0001
Bryophyte presence/absence	1	2.2	0.14	1	2.4	0.12
Warmed x shrub	1	0.1	0.78	1	2.8	0.10
Warmed x graminoid	1	0.4	0.55	1	0.6	0.44
Warmed x bryophyte	1	0.1	0.70	1	1.2	0.28
Shrub x graminoid	1	8.5	0.005	1	13.9	0.0003
Shrub x bryophyte	1	0.0	0.92	1	2.4	0.12
Graminoid x bryophyte	1	1.3	0.26	1	10.5	0.002
(no significant 3 way interactions)						
<b><i>N<sub>2</sub>O (mg m<sup>-2</sup> hr<sup>-1</sup>)</i></b>			(n = 164)			(n = 167)
Warming	1	0.0	0.92	1	2.5	0.12
Shrub presence/absence	1	0.1	0.82	1	0.5	0.50
Graminoid presence/absence	1	0.7	0.39	1	1.3	0.27
Bryophyte presence/absence	1	0.2	0.65	1	1.1	0.29
Warmed x shrub	1	0.1	0.76	1	2.2	0.14
Warmed x graminoid	1	1.1	0.31	1	0.2	0.63
Warmed x bryophyte	1	0.4	0.55	1	1.4	0.25
Shrub x graminoid	1	1.0	0.32	1	0.8	0.37
Shrub x bryophyte	1	4.2	0.04	1	0.4	0.51
Graminoid x bryophyte	1	0.1	0.71	1	0.2	0.63
(no significant 3 way interactions)						



**Table 2.** DOC and DON in soil solution, and microbial biomass C and N in soils, sampled during the growing season. Values are means  $\pm$  s.e.

<b>Vegetation type</b>	<b>DOC</b> ( $\mu\text{g C g dry wt soil}^{-1}$ )	<b>DON</b> ( $\mu\text{g N g dry wt soil}^{-1}$ )	<b>Microbial Biomass C</b> ( $\text{mg C g dry wt soil}^{-1}$ )	<b>Microbial Biomass N</b> ( $\text{mg N g dry wt soil}^{-1}$ )
<b><i>Non-warmed</i></b>				
Control	1304 ( $\pm 188$ )	686 ( $\pm 99$ )	19.3 ( $\pm 2.7$ )	3.2 ( $\pm 0.2$ )
Shrub only	1514 ( $\pm 131$ )	814 ( $\pm 76$ )	11.8 ( $\pm 3.4$ )	2.6 ( $\pm 0.6$ )
Graminoid only	2135 ( $\pm 237$ )	1164 ( $\pm 199$ )	13.6 ( $\pm 2.6$ )	3.5 ( $\pm 0.6$ )
Bryophyte only	2115 ( $\pm 205$ )	1185 ( $\pm 128$ )	14.7 ( $\pm 2.6$ )	3.3 ( $\pm 0.5$ )
Shrub + Graminoid	1357 ( $\pm 179$ )	726 ( $\pm 86$ )	17.0 ( $\pm 1.3$ )	3.0 ( $\pm 0.2$ )
Shrub + Bryophyte	1928 ( $\pm 170$ )	1018 ( $\pm 79$ )	18.3 ( $\pm 2.2$ )	3.7 ( $\pm 0.5$ )
Graminoid + Bryophyte	2141 ( $\pm 161$ )	1128 ( $\pm 84$ )	13.5 ( $\pm 1.9$ )	2.5 ( $\pm 0.2$ )
No vegetation	2020 ( $\pm 306$ )	1065 ( $\pm 157$ )	16.8 ( $\pm 1.6$ )	3.4 ( $\pm 0.5$ )
<b><i>Warmed</i></b>				
Control	1901 ( $\pm 267$ )	1169 ( $\pm 235$ )	16.0 ( $\pm 0.9$ )	3.1 ( $\pm 0.2$ )
Shrub only	2331 ( $\pm 166$ )	1216 ( $\pm 98$ )	10.7 ( $\pm 1.9$ )	2.2 ( $\pm 0.7$ )
Graminoid only	2096 ( $\pm 149$ )	1108 ( $\pm 81$ )	11.7 ( $\pm 0.8$ )	3.0 ( $\pm 0.4$ )
Bryophyte only	1900 ( $\pm 56$ )	1026 ( $\pm 31$ )	9.2 ( $\pm 1.8$ )	2.1 ( $\pm 0.6$ )
Shrub + Graminoid	1722 ( $\pm 240$ )	951 ( $\pm 123$ )	16.8 ( $\pm 1.0$ )	3.0 ( $\pm 0.8$ )
Shrub + Bryophyte	1531 ( $\pm 104$ )	789 ( $\pm 51$ )	15.8 ( $\pm 3.2$ )	3.6 ( $\pm 1.3$ )
Graminoid + Bryophyte	2376 ( $\pm 134$ )	1302 ( $\pm 83$ )	17.2 ( $\pm 0.8$ )	3.7 ( $\pm 0.5$ )
No vegetation	2502 ( $\pm 336$ )	1363 ( $\pm 145$ )	13.7 ( $\pm 0.8$ )	3.6 ( $\pm 0.4$ )



**FIGURE LEGENDS****Figure 1. Net ecosystem CO<sub>2</sub> exchange from the plant manipulation and warming****experiment.** Data are means (mg CO<sub>2</sub> m<sup>-2</sup> hr<sup>-1</sup>) for all sampling dates +/- standard error.

White bars are for non-warmed and black bars are for warmed experimental field plots. Data

are split between: growing season of May to September (left), and non-growing season of

October to April (right). Negative values represent a net sink and positive values represent a

net source for CO<sub>2</sub>.**Figure 2. Ecosystem respiration from the plant manipulation and warming experiment.**Data are means (mg CO<sub>2</sub> m<sup>-2</sup> hr<sup>-1</sup>) for all sampling dates +/- standard error. White bars are

for non-warmed and black bars are for warmed experimental field plots. Data are split

between: growing season of May to September (left), and non-growing season of October to

April (right).

**Figure 3. Methane flux from the plant manipulation and warming experiment.**Data are means (mg CH<sub>4</sub> m<sup>-2</sup> hr<sup>-1</sup>) for all sampling dates +/- standard error. White bars are for non-

warmed and black bars are for warmed experimental field plots. Data are split between:

growing season of May to September (left), and non-growing season of October to April

(right).

**Figure 4. Nitrous oxide flux from the plant manipulation and warming experiment.**Data are means (mg N<sub>2</sub>O m<sup>-2</sup> hr<sup>-1</sup>) for all sampling dates +/- standard error. White bars are

for non-warmed and black bars are for warmed experimental field plots. Data are split

between: growing season of May to September (left), and non-growing season of October to

April (right).







